

Workshop on Objective Seasonal Forecast (OSF) production over MedCOF region. 2-4 July 2024 Madrid, Spain

POTENTIAL IMPROVEMENTS THROUGH CALIBRATION AND DOWNSCALING: BLENDING ANALOGS AND MLR

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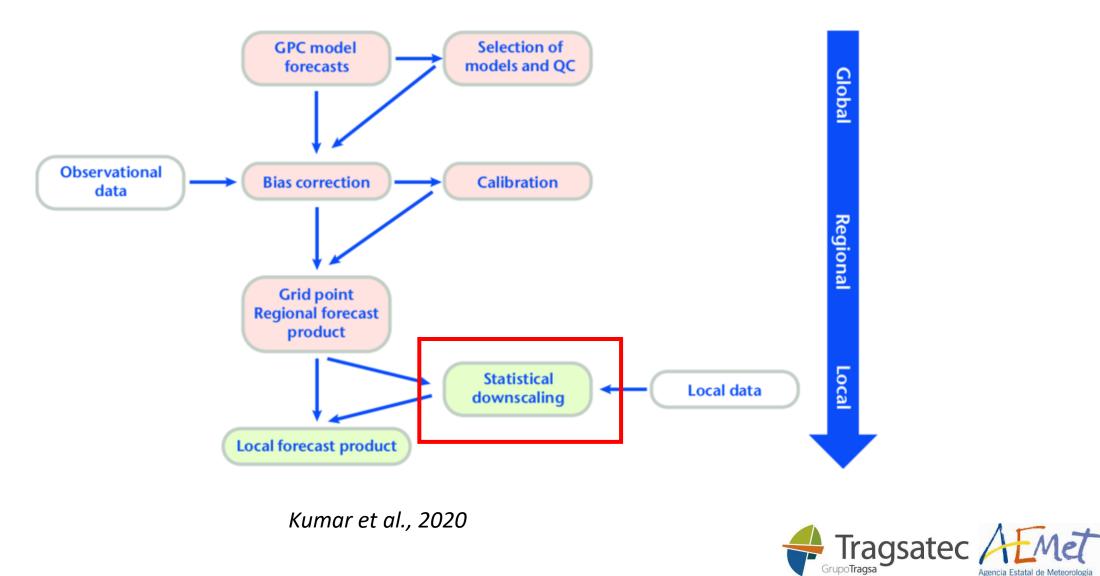


INDICE

- Context of the downscaling process: important aspects
- Downscaling AEMET algorithms: analogs, multivariable linear regression (MLR) and weather types (significant predictors)
- ERAInterim Validation
- Case example: SEAS5.1 Forecast last extended winter (NDJFM 2023-2024): added value of downscaling
 - Anomaly compared to the Hindcast (1997-2016)
 - Probability of terciles
 - Verification compared to ERA5
 - Skill (ROCArea) of upper and lower terciles
- Conclusions
- References



Outline of the recommended procedure to develop an objective seasonal forecast

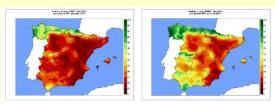


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Important aspects of downscaling process

Select representative predictors from Seasonal Forecast Systems (low resolution, e.g. SEAS5.1 10) ERA Interim 2-meter Temperature, mean of 1989 -73 -68 -63 -58 -53

Have **quality observational dataset** (high resolution, e.g. 5km)

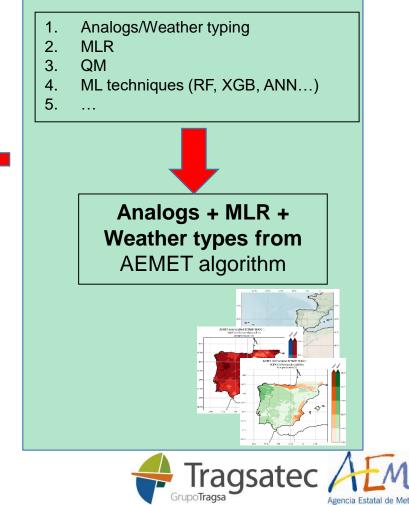


Serie de precipitación diaria en rejilla con fines climáticos

Nota técnica 24 de AEMET

Candelas Peral García Beatriz Navascués Fernández-Victorio Petra Ramos Calzado

Select statistical downscaling methods



Analogs downscaling algorithm from AEMET

(Petisco et al. (2008a), Petisco et al. (2008b), Amblar et al. (2017))

What it does: Downscaling of precipitation and maximum and minimum temperature of the seasonal forecast on a grid. It uses analogous synoptic situations in a past reference period based on Euclidean distances. Also uses regression and weather types.

Input data:

- Reanalysis (ERA Interim, ERA 5,...) and seasonal forecasting (Hindcast and Forecast (SEAS5.1))

- Observational High resolution data (ROCIO_IBEB 5km AEMET grid).

Output data:

- Temperature and precipitation data on a 5Km grid

Domain: centered on IP



ANALOGS + MLR downscaling algorithm from AEMET: TEMPERATURE

CALIBRATION and VALIDATION PERIOD (1981-1996) REANALYSIS + OBSERVATIONS

1. Obtain a set of analogs associated with an estimation day

DATA

- ERAInterim (1.5° x 1.5°)
- Geostrophic wind components
 1000-500 hPa from geopotential (12h)
- Domain: wider than IP

2. Regression model (with a set of selected analogs)

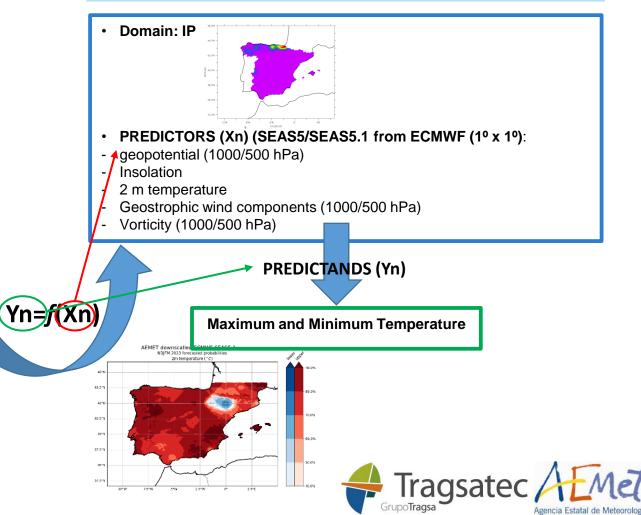
PREDICTANDS (Yn) (ROCIO_IBEB 5km):

- Tmax y Tmin

PREDICTORS (Xn) (ERAInterim 1.5°):

- geopotential (1000/500 hPa)
- Insolation
- 2 m temperature
- Geostrophic wind components (1000/500 hPa)
- Vorticity (1000/500 hPa)

ESTIMATION PERIOD (Hindcast (1997-2016) Forecast (2023-2024)) SEASONAL FORECAST SYSTEMS



WEATHER TYPES + ANALOGS downscaling algorithm from AEMET: PRECIPITATION

Yn=f(Xn)

CALIBRATION and VALIDATION PERIOD (1981-1996) REANALYSIS + OBSERVATIONS

- 1. Weather types: seeds (Lund, 1963) + k-means (Anderberg, 1973)
- 2. Synoptic analogs for each weather type

DATA

- ERAInterim (1.5° x 1.5°)
- Geostrophic wind components 1000-500 hPa from geopotential (12h)
- Domain: wider than IP

3. Stepwise regression: Significance of each predictor in a linear regression model, for each weather type

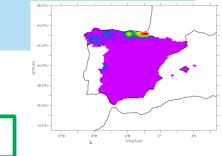
PREDICTAND (ROCIO_IBEB 5km): Precipitation Potential PREDICTORS:

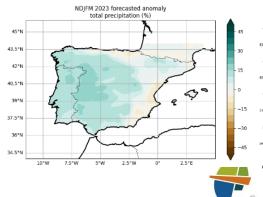
- geopotential (1000,500hPa)
- Relative humidity (700hPa)
- Relative vorticity (1000hPa)
- Geostrophic wind components (1000, 500hPa)
- Temperature (1000, 850, 700, 500hPa)

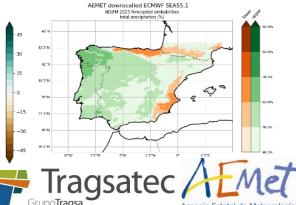
ESTIMATION PERIOD (Hindcast (1997-2016) // Forecast (2023-2024)) SEASONAL FORECAST SYSTEMS

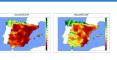
Precipitation

- 1. Obtain synoptic analogs
- 2. Determine to which synoptic type it belongs -> significant predictors
- 3. Local similarity using significant predictors





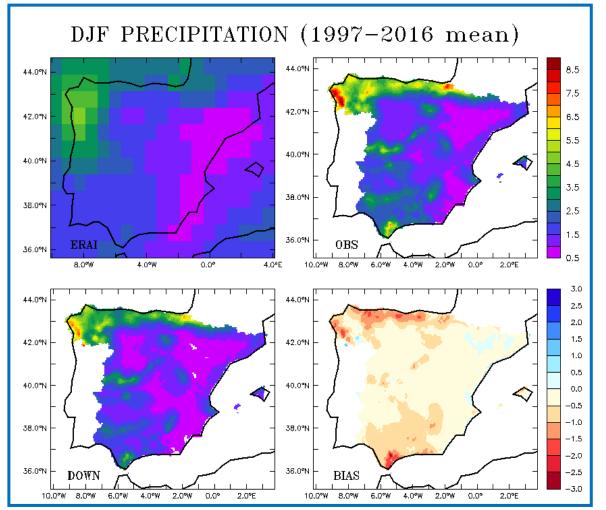




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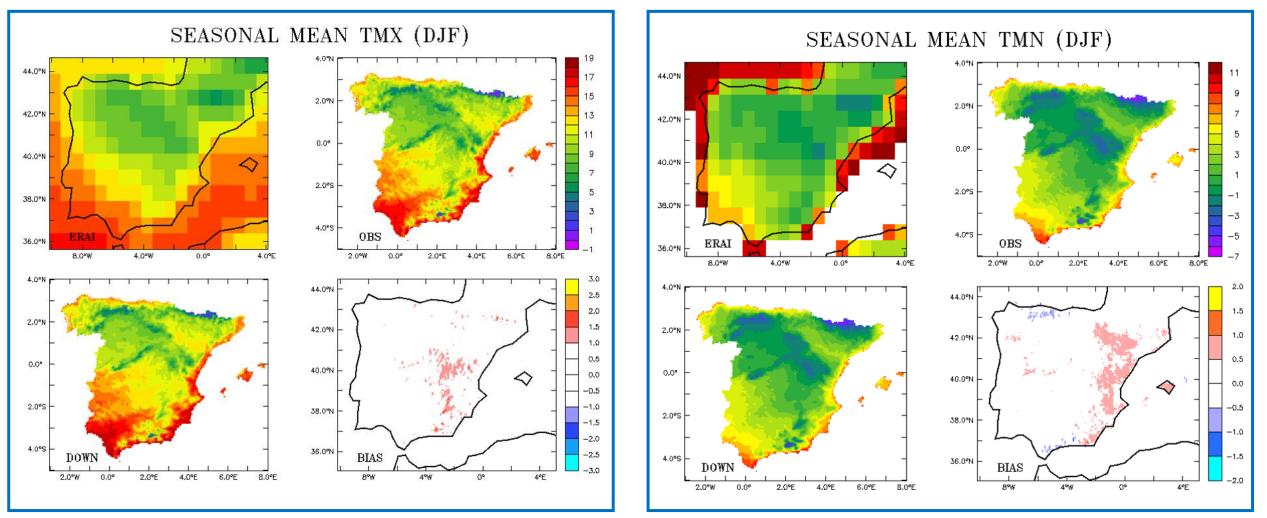
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PRECIPITATION ERAInterim evaluation

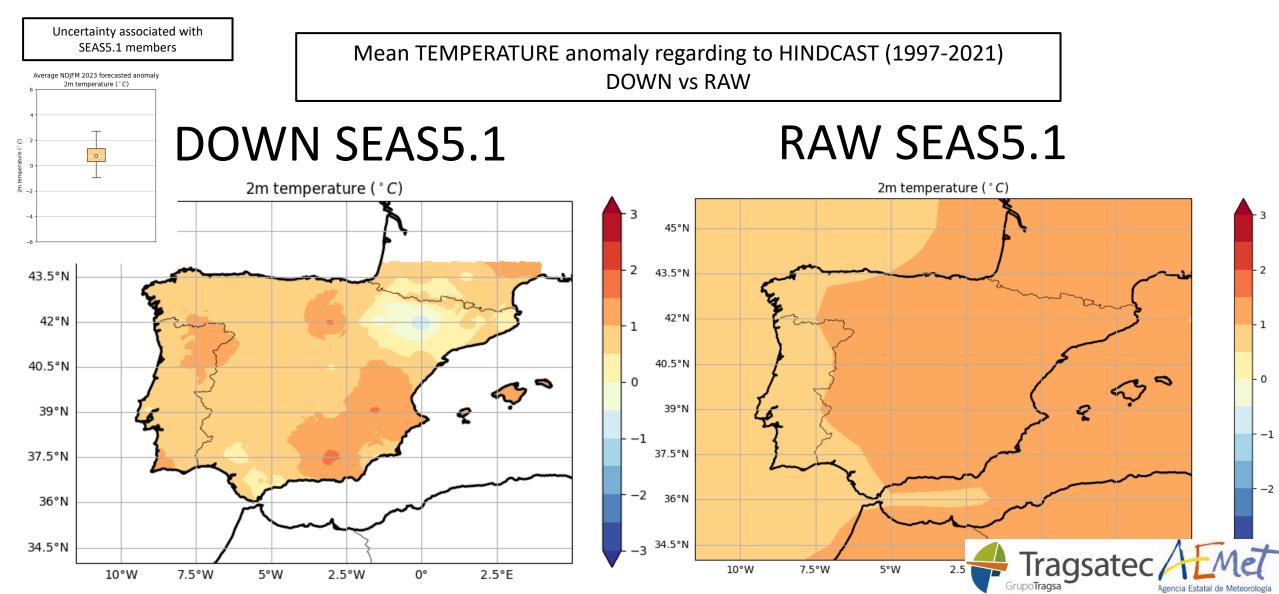


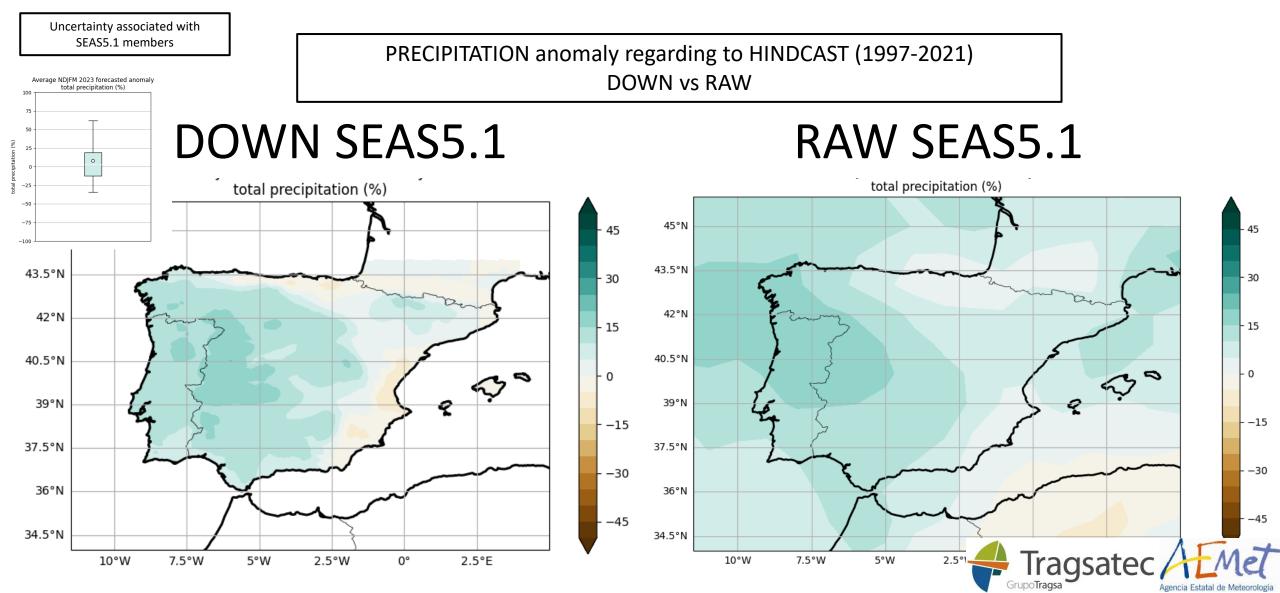
 Example of the spatial distribution of the downscaling daily winter precipitation (mm/day) of ERA Interim (ERAI, upper-left panel), AEMET observational database (OBS, upper-right panel), downscaling ERA Interim-based (DOWN, lower-left panel) and of the bias (BIAS (DOWN-OBS), lower-left panel), average over the period 1997-2016.

MAXIMUM and MINIMUM TEMPERATURE ERAInt evaluation



Example of the spatial distribution of the downscaling daily winter maximum and minimum temperature (°C) of ERA Interim (ERAI, upper-left panel), AEMET observational database (OBS, upper-right panel), downscaling ERA Interim-based (DOWN, lower-left panel) and of the bias (BIAS (DOWN-OBS), lower-left panel), average over the period 1997-2016.

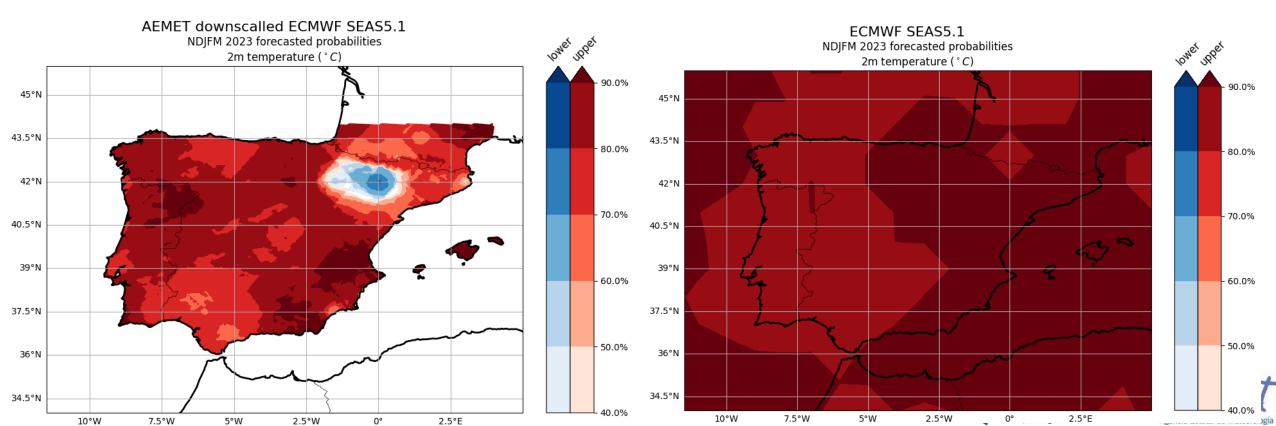




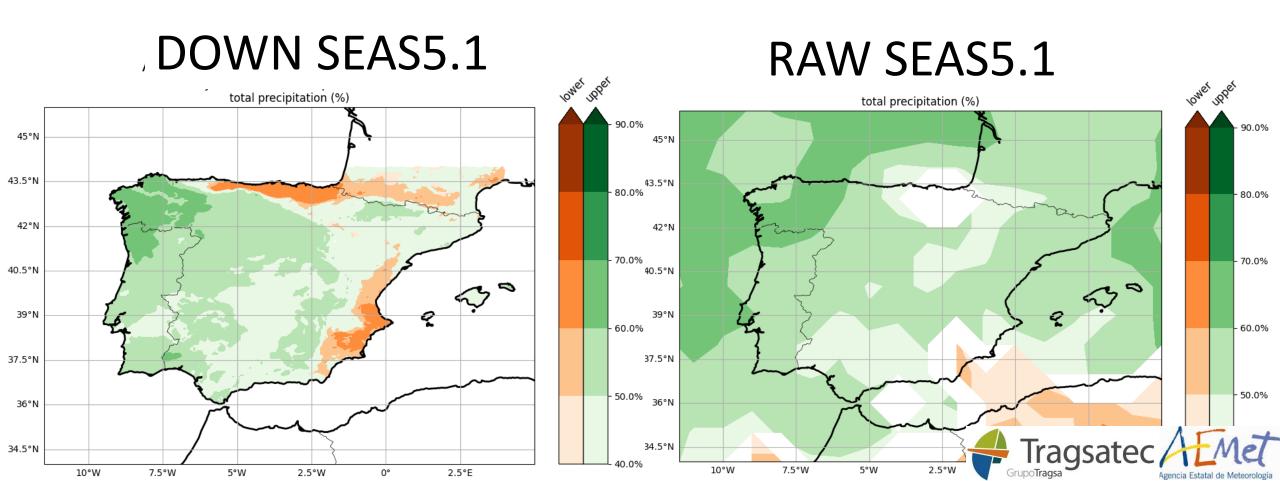
Most likely upper and lower terciles probability for mean TEMPERATURE DOWN vs RAW

DOWN SEAS5.1

RAW SEAS5.1



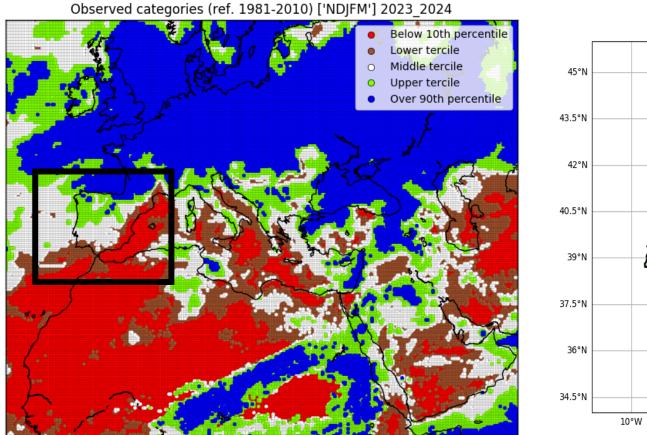
Upper/Lower terciles Probability of occurrence for PRECIPITATION DOWN vs RAW

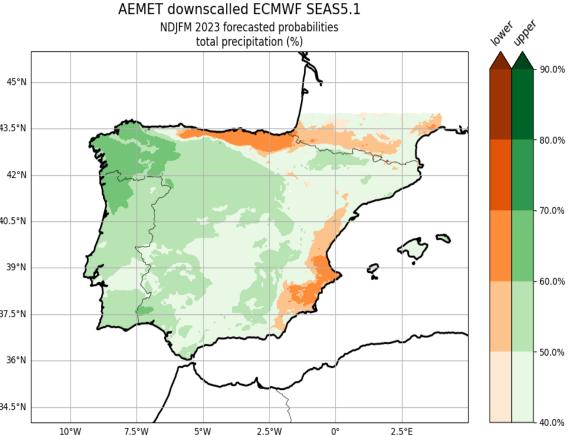


NDJFM 2023-2024 verification

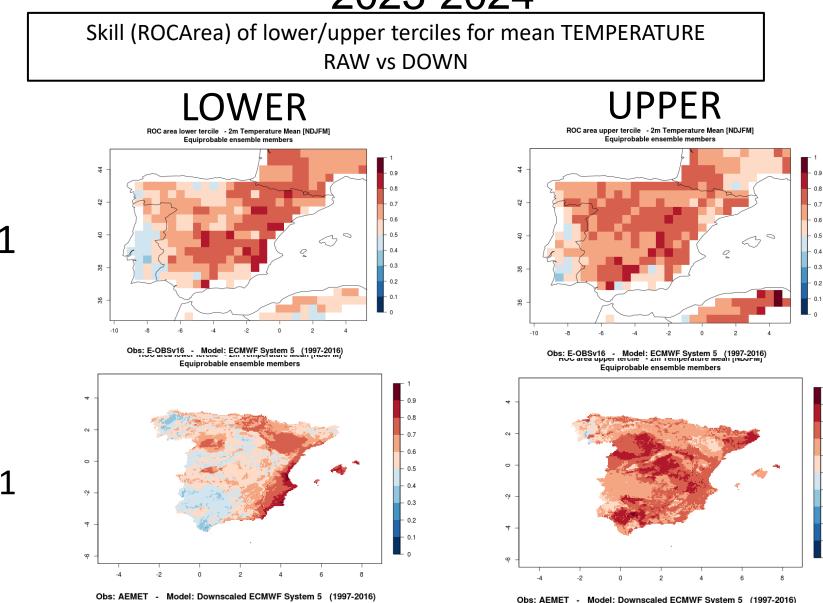
ERA5

DOWN SEAS5.1









0.9

8.0

0.7

0.6

0.5

0.4

0.3

0.2

0.1

sate

RAW SEAS5.1

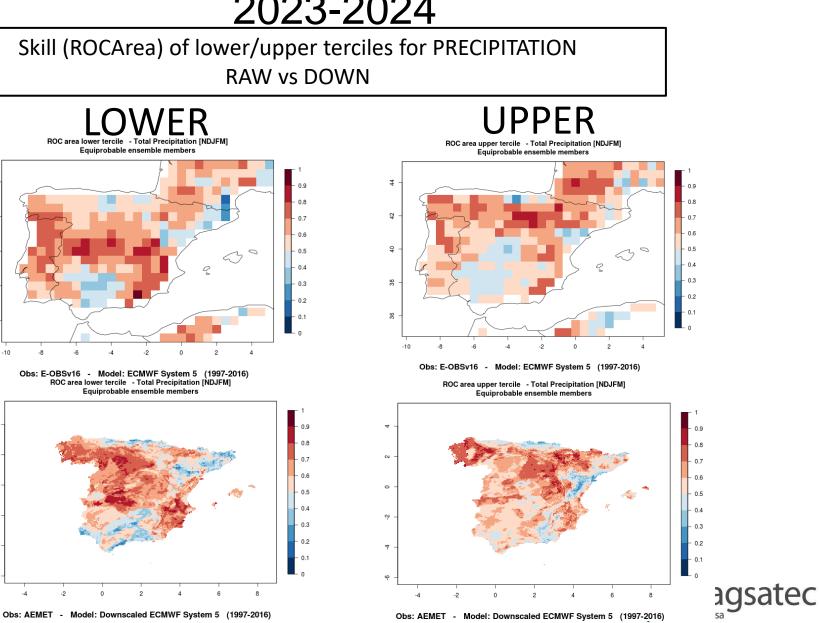
DOWN SEAS5.1

RAW SEAS5.1

4

얶

DOWN SEAS5.1



Conclusions

- It is important to take account representative predictors from Seasonal Forecast Systems, quality observational dataset and a proper downscaling method.
- The evaluation (1997-2016) of precipitation and maximum and minimum temperature estimated by the AEMET analog method shows that the methodology can provide geographical detail compared to ERAInterim, showing a very accurate spatial pattern concerning the AEMET observational dataset.
- During the extended winter, for both precipitation and mean temperature, the downscaling provides added value compared to the raw SEAS5.1 data in the following points:
 - The increased resolution of the seasonal forecast allows it to be used by resource management models in sectors such as hydrology or agriculture.
 - The greater spatial detail of the probability of occurrence allows a more realistic prediction to be made.
 - The seasonal prediction skill (ROC area) is improved both spatially and in terms of the number of occurrences.



References

- Amblar, M. P., Casado Calle, M. J., Pastor Saavedra, M. A., Ramos Calzado, P., and Rodríguez Camino, E. (2017). Guía de escenarios regionalizados de cambio climático sobre España a partir de los resultados del IPCC-AR5.
- Kumar, Arun & Ceron, Jean-Pierre & Coelho, Caio & Ferranti, L. & Graham, Richard & Jones, David & Merryfield, William & Muñoz, Ángel & Pai, Sivananda & Rodriguez-Camino, Ernesto. (2020). WMO Guidance on Operational Practices for Objective Seasonal Forecasting.
- Petisco de Lara, S. (2008a). Método de regionalización de precipitación basado en análogos: Explicación y validación. AEMET, Madrid.
- Petisco de Lara, S. (2008b). Método de regionalización de temperaturas basado en análogos: Explicación y validación. AEMET, Madrid.



Thank you for your attention!



La distancia euclídea entre los <u>campos tipificados</u> de cada componente *c* de dos elementos *m* y *n*, donde *c* representa una componente cualquiera $(u_{1000}, v_{1000}, u_{500}, v_{500})$ del viento geostrófico, vendría dada por la expresión [1].

$$d_{c} = \sqrt{\frac{\sum_{i=1}^{n} W_{i} (c_{mi} - c_{ni})^{2}}{\Sigma W_{i}}}$$
[1]

en donde:

 c_{mi} y c_{ni} representan el valor de la componente c tipificada de viento geostrófico en el punto de rejilla *i* para dos situaciones sinópticas *m*, *n* cualesquiera cuya distancia se quiere calcular;

w, es la ponderación correspondiente al punto de rejilla *i*;

n = 198 es el número total de puntos de rejilla del campo.

Una vez calculadas las cuatro distancias d_c correspondientes a cada uno de los cuatro campos de componentes de viento geostrófico, se toma la media de ellas como distancia D_s entre los elementos considerados.

Analogs downscaling algorithm from AEMET

Depending on variable, methodology is different:

Precipitation: reanalysis synoptic regimes are classified according to weather types (k-means, [Anderberg1973]) and significant predictors are determined for each of them. Finally, the group to which our problem day belongs is assigned, and rainfall is estimated by regression with the significant predictors for its weather type.

Tmax and Tmin: a multiple linear regression is carried out between some predictor variables of the reanalysis and the temperature data of the observational grid. Once the adjustment has been made, the temperature estimate is made by substituting, in the adjusted equation, the values that the predictor variables take on the problem day of the model.



